Miocene Planktonic Foraminiferal Biozonation for South Sumatra Basin, Indonesia

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Abstract

Planktonic foraminifera are widely utilized for marine sediments analysis and recent developments in the geochronology of planktonic foraminifera has improved their values as a tool for sedimentary basin analysis.

Assessment on the applicability of the "standard" foraminiferal biozonations (e.g., Bolli, 1966; Blow, 1969; and Postuma, 1971) to the biostratigraphy data of the South Sumatra Basin suggests that the standard zones, which are mostly interval-zones defined by first occurrence events, are hardly applicable to the biostratigraphy data derived from the examination of well-cuttings routinely and continuously used in most sedimentary basin studies in Indonesia.

The proposed biozonation in this paper were constructed based on two considerations, i.e. information on the biostratigraphy events as indicated in regional- and global-scale biozonation schemes and the occurrence of foraminifera species in the South Sumatra Basin.

The correlation between the proposed zonation and that of Blow's (1969) scheme suggests that the proposed zonation is basically a Blow's zonation and, accordingly, may be perceived as a modified Blow's zonation for South Sumatra Basin.

Keywords: planktonic foraminifera, biostratigraphy zonation; South Sumatra Basin

INTRODUCTION

Planktonic foraminiferal zonation scheme is one of the most important tools for the study of sedimentary basins. Fossils represented in the zonation schemes provide practical age control that is critical for understanding of basin evolution and are useful for correlation of marine sedimentary basin fills.

The planktonic foraminiferal zonation that have been used most in Indonesia basins studies is that of Blow, originally established almost 50 years ago in 1969. Most of the zones in the Blow zonation, however, are interval-zones defined by first occurrence events that are hardly applicable to the biostratigraphy data derived from the examination of well-cuttings routinely and continuously used in most sedimentary basin studies in Indonesia.

This paper proposes a Miocene planktonic foraminiferal biozonation for South Sumatra Basin that is more applicable for subsurface basin studies than most of "standard" zonation as it gives special emphasizes to the last appearance datum (LAD) of every species that may have bio chronostratigraphic

significance. It may be perceived as a modified form of the Blow zonation (1969), designed specifically to accommodate subsurface biostratigraphy information derived from the study of borehole cuttings and the types of Miocene planktonic foraminifera observable in the South Sumatra Basin sediments.

GEOLOGIC SETTING

Largely onshore, the South Sumatra Basin is situated on the southern end of the island of Sumatra (figure 1). It is partially separated from the Central Sumatra Basin to the north by a large exposure of crystalline and metasedimentary basement (i.e., the Tigapuluh Mountains) and from the Sunda Basin in the Java Sea to the southeast by the Lampung High and its northward extension in the islands of Bangka and Belitung. To the northeast, the South Sumatra Basin fills thin out over the Sunda-land basement, while its west and southwest side is limited along the margins of the Barisan Mountains.

The South Sumatra Basin comprises a series of semi-connected NNW-SSE

trending depressions, formed during the early Tertiary rifting phase (Adiwidjaja and De Coster, 1973; De Coster, 1974; Suhendan, 1984; Daly et al., 1987; 1985; Pertamina-Beicip, Pertamina BPPKA, 1997; Barber and Crow, 2005; Barber et al., 2005; De Smet and Barber, (figure 1). The depositional fill of the South Sumatra Basin began when the aforementioned depressions were filled with locally sourced continental volcaniclastic sediments of the Lahat Formation (Suhendan, 1984; Pertamina-Beicip, 1985; Barber et al., 2005; De Smet and Barber, 2005) (figure 2).

When the structuring associated with the rifting phase waned during the late Oligocene time, the early evolutionary stage of the South Sumatra Basin was marked by transgression as a result of thermal saq and eustatic gain (Netherwood, 2000; Barber et al., 2005). sediments of the Talangakar Formation were deposited then in basinal lows. The Talangakar was deposited in a diverse range of environments and it consists of interbedded sandstones, shales, and coals.

With continued transgression into the early Miocene, the Baturaja Formation developed on structural highs as lowrelief. low-energy, carbonate-muddominated banks (Situmeang et al., 1992; Longman et al., 1992). The Gumai Formation, which consists almost exclusively of fine-grained siliciclastic rocks with thinly bedded sandstone and limestone intercalations, were deposited off-bank in deeper water. During the middle Miocene maximum transgression, the Gumai seal across the region, including those formerly covered by carbonate buildups (De Coster, 1974).

During middle Miocene, the variations in the angle and rate of convergence in the Sumatran subduction system leading to extension or compression in the back arc basinal areas (Barber et al., 2005). The event was manifested by the uplift of the Barisan Mountains, the increasing volcanic activities in the forearc region, and the formation of anticlines and faulted anticlines in the South Sumatra Basin (Netherwood, 2000; Barber et al.,

2005). A middle Miocene – Pliocene regressive phase of deposition commenced with the shallow-marine to deltaic Air Benakat and Muaraenim Formations, which consists mainly of sandstone and fine-grained siliciclastic rocks with coal bed intercalations in the Muara Enim Formation.

As compression and volcanism continued, sedimentation from the Pliocene onward were dominated by the increasing volcanic materials from the west and southwest of the basin, leading to the deposition of the Kasai Formation (Pleistocene) that rests with unconformity the Muaraenim on Formation. The Kasai is composed of conglomerates, tuffaceous sandstones, and tuffs with lignite and silicified wood.

DATA AND METHODS

Raw data consulted for this paper the borehole occurrences planktonic foraminifera species in the subsurface South Sumatra Basin—were retrieved from paleontology biostratigraphy reports of a total 30 exploration wells. Original paleontology biostratigraphy analyses were conducted by various service companies from 1971 to 2001.

The bulk of the aforementioned biostratigraphy data derived from ditch cutting samples. Interval of analyzed samples were not necessarily the same from one well to another; they were ranging from about 3 m to about 100 m. Sidewall cores of selected horizon were also analyzed in almost all studied wells. "standard" Several planktonic foraminiferal zonation schemes—those proposed by Bolli (1966), Blow (1969), Postuma (1971), Srinivasan and Kenneth (1981), Bolli and Saunders (1985)—as well as regional zonation schemes (Pringgoprawiro et al., 1978; Ikebe and Chiji, 1978; Van Gorsel, 1988; Kadar et al., 1996; Van Gorsel et al., 2014) were assessed for their applicability to the available biostratigraphy data from the South Sumatra Basin. At the same time, the stratigraphic range of foraminifera observable in Miocene sediments of the South Sumatra Basin was revised to accommodate the results of recent

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studies (e.g., Berggren et al., 1995; Wade et al., 2011; Li et al., 2015; BouDagher-Fadel, 2015). The results of the assessment were used to construct a

planktonic foraminiferal zonation for the South Sumatra Basin (figure 3), using the biostratigraphy zonation nomenclature of Wade *et al.* (2011).

DEF	DEFINITIONS OF PLANKTONIC FORAMINIFERAL ZONES			
		Globorotalia kugleri Zone (N4)		
Definition	:	Biostratigraphic interval between the FAD of <i>Globorotalia</i> kugleri Bolli and the LAD of <i>Globorotalia</i> kugleri Bolli.		
Category	:	Taxon range zone (TRZ).		
Estimated age	:	22.96-21.12 Ma (Gradstein <i>et al.</i> , 2004; Chunlian <i>et al.</i> , 2012; Li <i>et al.</i> , 2015), Early Miocene.		
Remarks	:	This zone is equivalent to the <i>Globorotalia kugleri</i> Zone of Bolli (1966), Postuma (1971), Gradstein <i>et al.</i> (2004), Li <i>et al.</i> (2015). <i>Globigerina ciperoensis</i> Bolli has its extinction near the base of this zone, at an estimated age of 22.90 Ma (Gradstein <i>et al.</i> , 2004a; Li <i>et al.</i> , 2015). As such, its first occurrence may be used as a proxy for the base of N4 zone.		

DEFINITIONS OF PLANKTONIC FORAMINIFERAL ZONES		
		Globoquadrina binaiensis Zone (N5)
Definition	:	Biostratigraphic interval between the LAD of <i>Globorotalia</i>
		kugleri Bolli and the LAD of Globoquadrina binaiensis Koch.
Category	:	Highest-occurrence zone (HOZ).
Estimated age	:	21.12-19.09 Ma (Gradstein <i>et al.</i> , 2004; Chunlian <i>et al.</i> , 2012;
		Li <i>et al.</i> , 2015), Early Miocene.
Remarks	:	The LAD of <i>Globoquadrina binaiensis</i> Koch, which is proposed to mark the top of this zone, has been known as one of those
		that mark biostratigraphic events in Indonesia (Pringgoprawiro et al., 1978; Van Gorsel, 1988; Kadar et al., 1996; Van Gorsel et al., 2014) and beyond (Pearson and Chaisson, 1997;
		Chunlian et al., 2012; Gradstein et al., 2004). A general practice of using the FAD of Globigerinita insueta Cushman and Stainforth to mark the top of N5 zone (e.g., Blow, 1969; Berggren et al., 1995), is not followed in this paper due to the scarcity of nominate species in the South Sumatra Basin.

DEFINITIONS OF PLANKTONIC FORAMINIFERAL ZONES			
		Catapsydrax dissimilis Zone (N6)	
Definition	:	Biostratigraphic interval between the LAD of <i>Globoquadrina</i> binaiensis Koch and the LAD of Catapsydrax dissimilis Cushman and Bermudez.	
Category	:	Highest-occurrence zone (HOZ).	
Estimated age	:	19.09-17.54 Ma (Gradstein <i>et al.</i> , 2004; Chunlian <i>et al.</i> , 2012), Early Miocene.	
Remarks	:	Catapsydrax unicavus Bolli, Loeblich and Tappan has its extinction at the top of this zone, at an estimated age of 17.54 Ma (Berggren et al., 1995; Gradstein et al., 2004; Li et al., 2015). As such, its LAD may be used as an alternative to mark the top of N6 zone.	

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DEFINITIONS OF PLANKTONIC FORAMINIFERAL ZONES		
		Globigerinatella insueta Zone (N7)
Definition	:	Biostratigraphic interval between the LAD of <i>Catapsydrax</i> dissimilis Cushman and Bermudez and the FAD of <i>Praeorbulina</i> sicana de Stefani.
Category	:	Partial range zone (PRZ).
Estimated age	:	17.54-16.40 Ma (Berggren <i>et al.</i> , 1995; Gradstein <i>et al.</i> , 2004; Wade <i>et al.</i> , 2011; Chunlian <i>et al.</i> , 2012), Early Miocene.
Remarks	:	Globigerinoides altiaperturus Bolli has its extinction at the top of this zone, at an estimated age of 16.40 Ma (Kadar et al., 1996; Chunlian et al., 2012). As such, its LAD may be used as an alternative to mark the top of N7 zone.

DEFINITIONS OF PLANKTONIC FORAMINIFERAL ZONES			
		Praeorbulina sicana Zone (N8)	
Definition	:	Biostratigraphic interval between the FAD of <i>Praeorbulina</i> sicana de Stefani and the FAD of <i>Orbulina</i> suturalis Brönnimann.	
Category	:	Partial range zone (PRZ).	
Estimated age	:	16.40-15.10 Ma (Berggren <i>et al.</i> , 1995; Gradstein <i>et al.</i> , 2004; Wade <i>et al.</i> , 2011), Middle Miocene.	
Remarks	:	According to Thompson and Abbott (2003), <i>Praeorbulina</i> glomerosa curva Blow and <i>Praeorbulina</i> glomerosa circularis Blow in Southeast Asia region have their extinction near the top of this zone, at an estimated age of 15.00 Ma. As such, their LAD may be used as proxies for the top of N8 zone.	

DEFINITIONS OF PLANKTONIC FORAMINIFERAL ZONES			
		Orbulina suturalis Zone (N9)	
Definition	•	Biostratigraphic interval between the FAD of <i>Orbulina suturalis</i> Brönnimann and the FAD of <i>Globorotalia peripheroacuta</i> Blow and Banner.	
Category	:	Lowest-occurrence zone (LOZ).	
Estimated age	:	15.10-14.23 Ma (Berggren <i>et al.</i> , 1995; Gradstein <i>et al.</i> , 2004; Wade <i>et al.</i> , 2011), Middle Miocene.	
Remarks	:	Praeorbulina transitoria Blow has its extinction near the top of this zone, at an estimated age of 14.24 Ma (Chunlian et al., 2012). As such, its LAD may be used as a proxy for the top of N9 zone.	

DEFINITIONS OF PLANKTONIC FORAMINIFERAL ZONES				
	Globorotalia peripheroacuta Zone (N10)			
Definition	:	Biostratigraphic interval between the FAD of <i>Globorotalia</i>		
		peripheroacuta Blow and Banner and the FAD of Globorotalia		
		praefohsi Blow and Banner.		
Category	:	Lowest-occurrence zone (LOZ).		
Estimated age	:	14.23-13.74 Ma (Wade <i>et al.</i> , 2011), Middle Miocene.		
Remarks	:	Globorotalia peripheroronda Blow has its extinction near the		
		top of this zone, at an estimated age of 13.77 Ma (Wade et al.,		
		2011). As such, its LAD may be used as a proxy for the top of		
		N10 zone.		

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DEFINITIONS OF PLANKTONIC FORAMINIFERAL ZONES			
		Globorotalia praefohsi Zone (N11)	
Definition	:	Biostratigraphic interval between the FAD of <i>Globorotalia</i>	
		praefohsi Blow and Banner and the FAD of Globorotalia fohsi	
		Cushman and Ellisor.	
Category	:	Lowest-occurrence zone (LOZ).	
Estimated age	:	13.74-13.34 Ma (Wade <i>et al.</i> , 2011), Middle Miocene.	
Remarks	:	The existence of this zone in South Sumatra Basin is	
		questionable as the present available data shows no record of	
		foraminifera having their extinction in this zone (see discussion	
		below).	

DEFINITIONS OF PLANKTONIC FORAMINIFERAL ZONES			
		Globorotalia fohsi Zone (N12)	
Definition	:	Biostratigraphic interval between the FAD of Globorotalia fohsi	
		Cushman and Ellisor and the LAD of Globorotalia fohsi	
		Cushman and Ellisor.	
Category	:	Taxon range zone (TRZ).	
Estimated age	:	13.34-11.75 Ma (Wade <i>et al.</i> , 2011), Middle Miocene.	
Remarks	:	The existence of this zone in South Sumatra Basin is	
		questionable as the present available data could not be used to	
		recognize the top of the Blow's (1969) N11 zone (see remarks	
		in Globorotalia praefohsi zone above and discussion below).	

DEFINITIONS OF PLANKTONIC FORAMINIFERAL ZONES			
		Globigerinoides ruber Zone (N13)	
Definition	:	Biostratigraphic interval between the LAD of <i>Globorotalia fohsi</i> Cushman and Ellisor and the FAD of <i>Globigerina nepenthes</i> Todd.	
Category	:	Partial range zone (PRZ).	
Estimated age	:	11.75-11.55 Ma (Wade <i>et al.</i> , 2011), Middle Miocene.	
Remarks	:	Globigerinoides subquadratus Brönnimann has its extinction near the top of this zone, at an estimated age of 11.46 Ma (Wade et al., 2011). As such, its LAD may be used as a proxy for the top of N13 zone.	

DEF	DEFINITIONS OF PLANKTONIC FORAMINIFERAL ZONES		
Globe	orot	alia nepenthes - Globorotalia mayeri Zone (N14)	
Definition	:	Biostratigraphic interval between the FAD of <i>Globigerina</i> nepenthes Todd and the LAD of <i>Globorotalia mayeri</i> Cushman and Ellisor.	
Category	:	Concurrent-range zone (CRZ).	
Estimated age	:	11.55-10.46 Ma (Wade <i>et al.</i> , 2011; Li <i>et al.</i> , 2015), Late Miocene.	
Remarks	:	Globorotalia siakensis LeRoy, at least in low latitude region, is presently regarded as the synonym for Globorotalia mayeri Cushman and Ellisor (see discussion by Bolli and Saunders, 1985). The LAD of Globorotalia siakensis LeRoy in South Sumatra Basin, therefore, may be used as an alternative to mark the top of N14 zone.	

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DEFINITIONS OF PLANKTONIC FORAMINIFERAL ZONES			
		Globorotalia continuosa Zone (N15)	
Definition	:	Biostratigraphic interval between the LAD of <i>Globorotalia</i>	
		mayeri Cushman and Ellisor or the LAD of Globorotalia	
		siakensis LeRoy and the FAD of Globorotalia acostaensis Blow.	
Category	:	Partial-range zone (PRZ).	
Estimated age	:	10.46-9.79 Ma (Wade <i>et al.</i> , 2011; Li <i>et al.</i> , 2015), Late	
		Miocene.	

DEFINITIONS OF PLANKTONIC FORAMINIFERAL ZONES											
Globorotalia acostaensis Zone (N16)											
Definition	Definition : Biostratigraphic interval between the FAD of <i>Globorotalia</i>										
		acostaensis Blow and the FAD of Globorotalia plesiotumida									
		Blow and Banner.									
Category	:	Lowest-occurrence zone (LOZ).									
Estimated age	:	9.79-8.52 Ma (Wade <i>et al.</i> , 2011; Li <i>et al.</i> , 2015), Late									
		Miocene.									

DEFINITIONS OF PLANKTONIC FORAMINIFERAL ZONES									
Globorotalia plesiotumida Zone (N17)									
Definition	:	Biostratigraphic interval between the FAD of <i>Globorotalia</i> plesiotumida Blow and Banner and the FAD of <i>Globorotalia</i> tumida Brady.							
Category	:	Lowest-occurrence zone (LOZ).							
Estimated age	:	8.52-5.51 Ma (Wade <i>et al.</i> , 2011), Late Miocene.							
Remarks	:	Globoquadrina dehiscens Chapman, Parr and Collins has its extinction near the top of this zone, at an estimated age of 5.80 Ma (Berggren et al., 1995). As such, its LAD may be used as a proxy for the top of N17 zone. According to Berggren et al. (1995), Globorotalia lenguaensis Bolli has its extinction at the top of N17 zone. As such, its LAD may be used as an alternative to mark the top of N17 zone.							

DEFINITIONS OF PLANKTONIC FORAMINIFERAL ZONES								
Globorotalia tumida Zone (N18)								
Definition	:	Biostratigraphic interval between the FAD of <i>Globorotalia</i>						
		tumida Brady and the FAD of Sphaeroidinella dehiscens						
Category	:	Lowest-occurrence zone (LOZ).						
Estimated age	:	5.51-5.48 Ma (Wade <i>et al.</i> , 2011), Late Miocene.						

DISCUSSION

As it is indicated clearly from previous section, the proposed zonation (figure 3) has been constructed based on regional to global biostratigraphic events that have proved as time-honored datums. At the same time, however, we present secondary datums that may serve as proxies for those primary datums, especially in subsurface studies when the majority of biostratigraphic data derived from ditch cutting samples.

The correlation between the proposed zonation and that of Blow's (1969) original scheme and those proposed by Berggren et al. (1995) and Wade et al. (2011) (figure 4) suggests that the proposed zonation is basically a Blow's zonation and, accordingly, may be perceived as a modified Blow's zonation for South Sumatra Basin. The general similarity between the proposed zonation and the regional biozonation scheme (Lunt, 2013, in Van Gorsel et al., 2013) indicates that the proposed zonation may

be used as a basis to correlate the results of biostratigraphic studies in the South Sumatra Basin with those of studies conducted in other Southeast Asia basinal areas

The existence of Blow's (1969)fohsi and Globorotalia Globorotalia praefohsi zones in South Sumatra Basin is questionable as the present available data could not be used to recognize them. The taxonomic problem pertaining to Fohsella lineage (see Berggren et al., 1995 and Wade et al., 2011 for discussion on that matter) may also contribute to the void recognition of the Globorotalia praefohsi zone in South Sumatra Basin. Both, however, may be combined to form a Globorotalia praefohsi - Globorotalia fohsi concurrent-range zone (N11-N12).

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REFERENCES

- Adiwidjaja, P., and G. L. de Coster. 1973. Pre-Tertiary paleotopography and related sedimentation in South Sumatra. *Proceedings Indon. Petrol. Assoc. 2nd Ann. Conv.*, pp. 89-103.
- Barber, A. J., and M. J. Crow. 2005. Structure and structural history. *In*: A. J. Barber, M. J. Crow, dan J. S. Milson (eds.) Sumatra: Geology, Resources and Tectonis Evolution. *Geol. Soc.* (London) Memoir 31, pp.175-233.
- Barber, A. J., M. J. Crow, and M. E. M. de Smet. 2005. Tectonic evolution. *In*: A. J. Barber, M. J. Crow, dan J. S. Milson

- (eds.) Sumatra: Geology, Resources and Tectonis Evolution. *Geol. Soc.* (London) Memoir 31, pp. 234-259.
- Berggren, W. A., D. V. Kent, C. C. Swisher III, and M-P. Aubry. 1995b. A revised Cenozoic geochronology and chronostratigraphy. *In*: W. A. Berggren, D. V. Kent, M-P. Aubry, dan J. Handerbol (eds.) Geochronology, Time Scales, and Global Statigraphic Correlation. *Soc. Econ. Paleont. Mineral. Spec. Publ.* 54, pp. 129-212.
- Bolli, H. M., and J. B. Saunders. 1985.
 Oligocene to Holocene low latitude planktic foraminifera. *In*: H. M. Bolli, J. B. Saunders, dan K. Perch-Nielsen (eds.) *Plankton Stratigraphy*. Cambridge: Cambridge University Press. pp. 155-262.
- BouDagher-Fadel, M. K. Biostratigraphical and Geological Significance of Planktonic Foraminifera. 2nd ed. UCL Press. 298 p.
- Chunlian, L., H. Yi, W. Jie, Q. Guoquan, Y. Tingting, X. Lianze, and Z. Suqing. 2012. Miocene-Pliocene planktonic foraminiferal biostratigraphy of the Pearl River mount basin, northern South China Sea. *Jour. Palaeogeography*, vol. 1, no. 1, pp. 43-56
- Daly, M. C., B. G. D. Hooper, and D. G. Smith. 1987. Tertiary plate tectonics and basin evolution in Indonesia. Proceedings Indon. Petrol. Assoc. 16th Ann. Conv., Vol. I, pp. 399-428.
- De Coster, G. L. 1974. The geology of the Central and South Sumatra Basins. Proceedings Indon. Petrol. Assoc. 3rd Ann. Conv., pp. 77-110.
- De Smet, M. E. M., and A. J. Barber. 2005. Tertiary stratigraphy. *In*: A. J. Barber, M. J. Crow, dan J. S. Milson (eds.) Sumatra: Geology, Resources and Tectonis Evolution. *Geol. Soc.* (London) Memoir 31, pp. 86-97.
- Gradstein, F. M., J. G. Ogg, and A. G. Smith (eds.) 2004. *A Geologic Time Scale 2004*. Cambridge: Cambridge University Press. 589 p.
- Ikebe, N., and M. Chiji. 1978. Neogene datum-planes of western Pacific, a proposal for discussion. *Geol. Research and Development Centre Spec. Publ.* 1, pp. 159-161.
- Kadar, A. P., D. W. Paterson, and Hudianto. 1996. Successful techniques

- and pitfalls in utilizing biostratigraphic data in structurally complex terrain: Vico Indonesia's Kutei Basin experience. *Proceedings Indon. Petrol. Assoc. 25th Ann. Conv.*, Vol. I, pp. 313-331.
- Li, C.-F., Lin, J., Kulhanek, D.K., Williams, T., Bao, R., Briais, A., Brown, E.A., Chen, Y., Clift, P.D., Colwell, F.S., Dadd, K.A., Ding, W., Hernández Almeida, I., Huang, X.-L., Hyun, S., Jiang, T., Koppers, A.A.P., Li, Q., Liu, C., Liu, Q., Liu, Z., Nagai, R.H., Peleo-Alampay, A., Su, X., Sun, Z., Tejada, M.L.G., Trinh, H.S., Yeh, Y.-C., Zhang, C., Zhang, F., Zhang, G.-L., and Zhao, X., 2015. Methods. In Li, C.-F., Lin, J., Kulhanek, D.K., and the Expedition 349 Scientists, Proceedings of the Integrated Ocean Discovery Program, 349: South China Sea Tectonics: College Station, TX (International Ocean Discovery Program). http://dx.doi.org/10.14379/iodp.proc. 349.102.2015.
- Longman, M. L., C. T. Siemers, and T. Siwindono. 1992. Characteristics of low relief carbonate mudbank reservoir rocks, Baturaja Formation (Lower Miocene), Air Serdang and Mandala Fields, South Sumatra Basin, Indonesia. *In*: C. T. Siemers, M. W. Longman, R. K. Park, and J. G. Kaldi (eds.) *IPA Core Workshop Notes 1*, pp. 9-1 to 9-29.
- Netherwood, R. 2000. The petroleum geology of Indonesia, overview of Indonesia's oil and gas industry. In: *Indonesia 2000 Reservoir Optimization Conference*, Jakarta, PT. Schlumberger Indonesia, pp. 174-227.
- Pertamina-Beicip. 1985. Hydrocarbon Potential of Western Indonesia. Jakarta: Pertamina. 293 p.
- Pertamina BPPKA. 1997. Petroleum Geology of Indonesian Basins: Principles, Methods, and Applications. Vol. X, South Sumatra Basins. Jakarta: Pertamina BPPKA. 91 p.
- Postuma, J. A. 1971. *Manual of Planktonic Foraminifera*. Amsterdam: Elsevier. 420 p.
- Pringgoprawiro, H., N. Soeharsono, and F. X. Soejanto. 1978. Subsurface Neogene planktonic foraminiferal biostratigraphy of North-West Java Basin. *Geol. Research and*

- Development Centre Spec. Publ. 1, pp. 125-136.
- Situmeang, S. P., C. W. Zeliff, and R. A. Lorenz. 1992. Characterization of low relief carbonate banks, Baturaja Formation, Ramba A and B pools, South Sumatra, Indonesia. *In*: C. T. Siemers, M. W. Longman, R. K. Park, and J. G. Kaldi (eds.) *IPA Core Workshop Notes* 1, pp. 8-1 to 8-10.
- Srinivasan, M. S., and J. P. Kennett. 1981. A review of Neogene planktonic foraminiferal biostratigraphy: Applications in the equatorial and South Pacific. *In*: J. E. Warme, R. G. Doeglas, and E. L. Winterer (eds.) The Deep Sea Drilling Project: A Decade of Progress. *Soc. Econ. Paleont. Mineral. Spec. Publ.* 32, pp. 395-432.
- Suhendan, A. R. 1990. South Sumatra exploration and production history. *In*: Courteney, S., P. Cockcroft, R. Lorentz, R. Miller, H. L. Ott, P. Prijosoesilo, A. R. Suhendan, dan A. W. R. Wright (eds.), *Indonesia: Oil and Gas Fields Atlas, Vol. 3 South Sumatra*. Jakarta: Indonesian Petroleum Association. pp. 2-8.
- Thompson, P. R., and W. H. Abbott. 2003. Chronostratigraphy and microfossilderived sea-level history of the Qiong Dong Nan and Ying Ge Hai Basins, South China Sea. *In*: H. C. Olson and R. M. Leckie (eds.) Micropaleontologic Proxies for Sea-Level Change and Stratigraphic Discontinuities. *Society for Sedimentary Geology Spec. Publ.* 75, pp. 97-117.
- Van Gorsel, J. T. 1988. Biostratigraphy in Indonesia: Methods, pitfalls and new directions. *Proceedings Indon. Petrol. Assoc. 17th Ann. Conv.*, Vol. I, pp. 275-300.
- Van Gorsel, J. T., P. Lunt, and R. Morley. 2014. Introduction to Cenozoic biostratigraphy of Indonesia-SE Asia. *Berita Sedimentologi*, No. 29, pp. 6-40.
- Wade, B. S., P. N. Pearson, W. A. Berggren, and H. Palike. 2011. Review and revision of Cenozoic tropical planktonic foraminiferal biostratigraphy and calibration to the geomagnetic polarity and astronomical time scale. *Earth Sci. Reviews*, vol. 104, pp. 111-142.

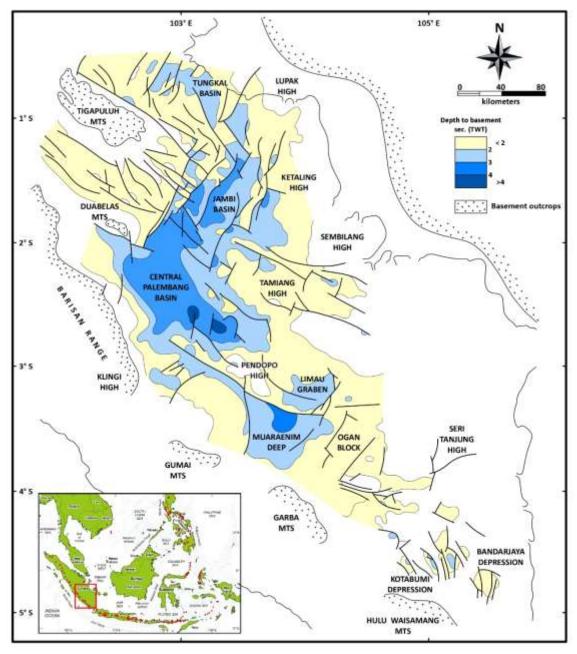


Figure 1. Geotectonic constellation and structural configuration of the South Sumatra Basin (modified from Pertamina-Beicip, 1985).

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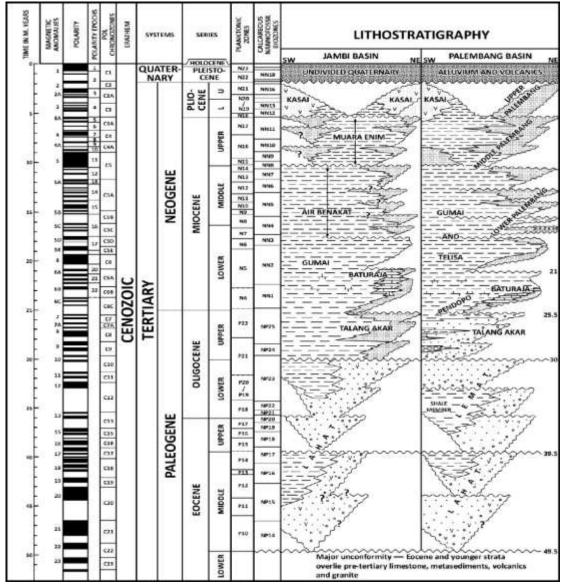


Figure 2. Lithostratigraphic scheme of the South Sumatra Basin (modified from Pertamina BPPKA, 1997).

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AGE Ma			PLA	NKTONIC FORAMINIFERAL BIOZONES	DATUMS FAD TLAD	SECONDARY DATUMS AS PROXIES				
PLIOCENE			N19							
	1	5.48	N18	Globorotalia tumida L-O-Z	oborotalia tumida					
		8.52	N17	Globorotalia plesiotumida L-O-Z	→ Globorotalia plesiotumida	◆¬ Globoquadrina defiscens (s.st)				
	¥	9.79	N16	Globorotalia acostaensis L-O-Z	Globorotalia acostaensis	Globorotalia siakensis				
		2020	N15	Globorotalia continuosa P-R-Z						
		10.46	N14	G. nepethes – G. mayeri C-R-Z	Globarotalia mayeri					
		11.55	N13	Globigerinoides ruber P-R-Z		Globigerinoides subquadratus (11.46)				
ш		11.75	N12	Globorotalia fohsi T-R-Z	Globorotalia fohsi					
MIOCENE		13.34	N11	Globorotalia praefohsi L-O-Z	Globorotalia praefahsi					
₩.		13.74	N10	Glaborotalia peripheroacuta L-O-Z		Globorotalia peripheroronda (23.77)				
		14,23	N9	Orbulina suturalis L-O-Z		Praeorbulina transitoria (14.74)				
		15.10	N8	Praearbulina sicana P-R-Z	J Orbulina suturalis	Praeorbulina giomerosa curva/circularis (15.00)				
	EARLY 5		N7	Globigerinatella insueta P-R-Z	- Praeorbulina sicana	Globigerinoides altiaperturus (25.40)				
		17.54	N6	Globorotalia dissimilis H-O-Z	Catopsydrox dissimilis Globoquadrina binalensis	Catopsydrax unicavus (17.54)				
		19.09	NS	Globorotalia binaiensis H-O-Z						
		21.12	N4	Globorotalia kugleri T-R-Z	Globorotalia kugleri	2447 V				
DUGO	DUGOCENE 22		P22	1, 2, 11, 2	a disconstanti kugien	Globigerina ciperaensis (22,90)				

Figure 3. Proposed Miocene planktonic foraminiferal biozonation for South Sumatra Basin.

Ma	AGE			Now (1969)	Berggren dik (1995b)		Wade dkk (2011)		Lunt (2013)		Present Study	
**	PLEISTS	CIM	N23	NP & Admirátio	MZS	AMF-12 (Nameson 13.17)	NZZ	MR S Denne (LPE)	N23 N22	N/ & treatments		
185	PLIOCENE	EARLY LATE	N21 N28 N38	10 6 contains	MZ1 MZ0 M19 M18	IAR Competito (158) IAR Competito (158) IAR Eddoor (178)	N21 N20 N19 N18	UR & Imperior (UR)	N21 N29		o curs	
11			N18	100 C Administration	MLT	NO E-works (ESS)	MLT	MA 4 (main (6.85)	NO.	SAID E distriction (b.ME)	MIT	Mt 5 bank (\$10)
		LATE	N17	100 2 ,000	MIS		N1.5	10 2 promote 2/80 10 2 sententi (24) 10 2 comi (189)	NUS	MI N NAME AND STREET A	M16 M15	ME 5 piceroste (676) ME 6 monace (676) (46 6 may 7164)
.200	ene	MIDDLE	N15 N34 N13	NO E-markets NO E-markets NO E-markets NO E-markets	M15 M14 N13 N22 M11	AND RESIDENCE STATE	NIA SNLSS NLS NLS NLS	Mar C. para (C. A)	NS4 NS3: NS2 NS3 NS3	AGE & programme (C. 16) AGE & programme (C. 16) AGE & programme (C. 16)	MG4 CMLEC MID MID MID	THE E-PART DESIGN
	MIDCENE	3	N12 N31 N30	MD 0,560 FIR 4 peoples MD 5 prophenous	NB NB	ME Experiment SAM	NR NR	MA E profession CASI NA E manufa (1918) NA F Appen (1948)	100	MO Complements (ILE) MI Selding (ILE) All Column (ILES)	**	MG E propression (SE) MG E values (SE) MG E values (SE)
		, o	MS MS	NE Colorb	NO NE	AM & south PASS	N7	MR C front (1735)	M7 M6	und calument prant	167	No. 6 amount 12 lbd
		EARLY	Mi	ME E-man	NS	No Physical (SSS)	15	MR Straggeronde (2008)	MS	NO SIGNATURE (FEE)	165	180 4 (1999) (212)
au	DUGOCENE	LATE	N3 (P22)	HP is principle:	MI MI	the digenostrates (Add	NA NO	Ma I hagen (SLTI)	N3 (P22)	*** ***********************************	NS. (P22)	MB il lugare (MIN)

Figure 4. Correlation between the proposed biozonation with some "standard" (Blow, 1969; Berggren *et al.*, 1995; Wade *et al.*, 2011) and regional (Lunt, 2013) biozonation schemes.